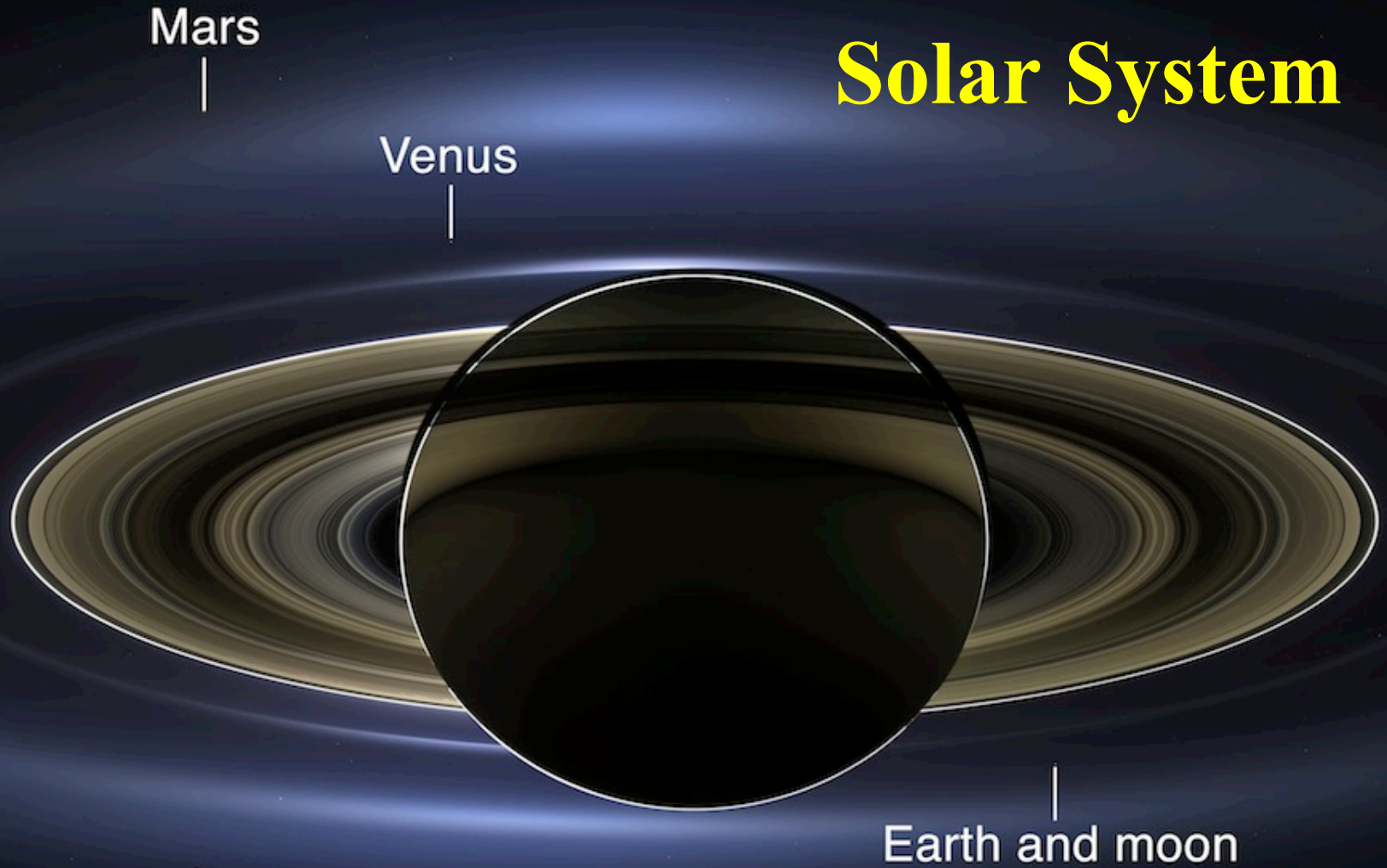


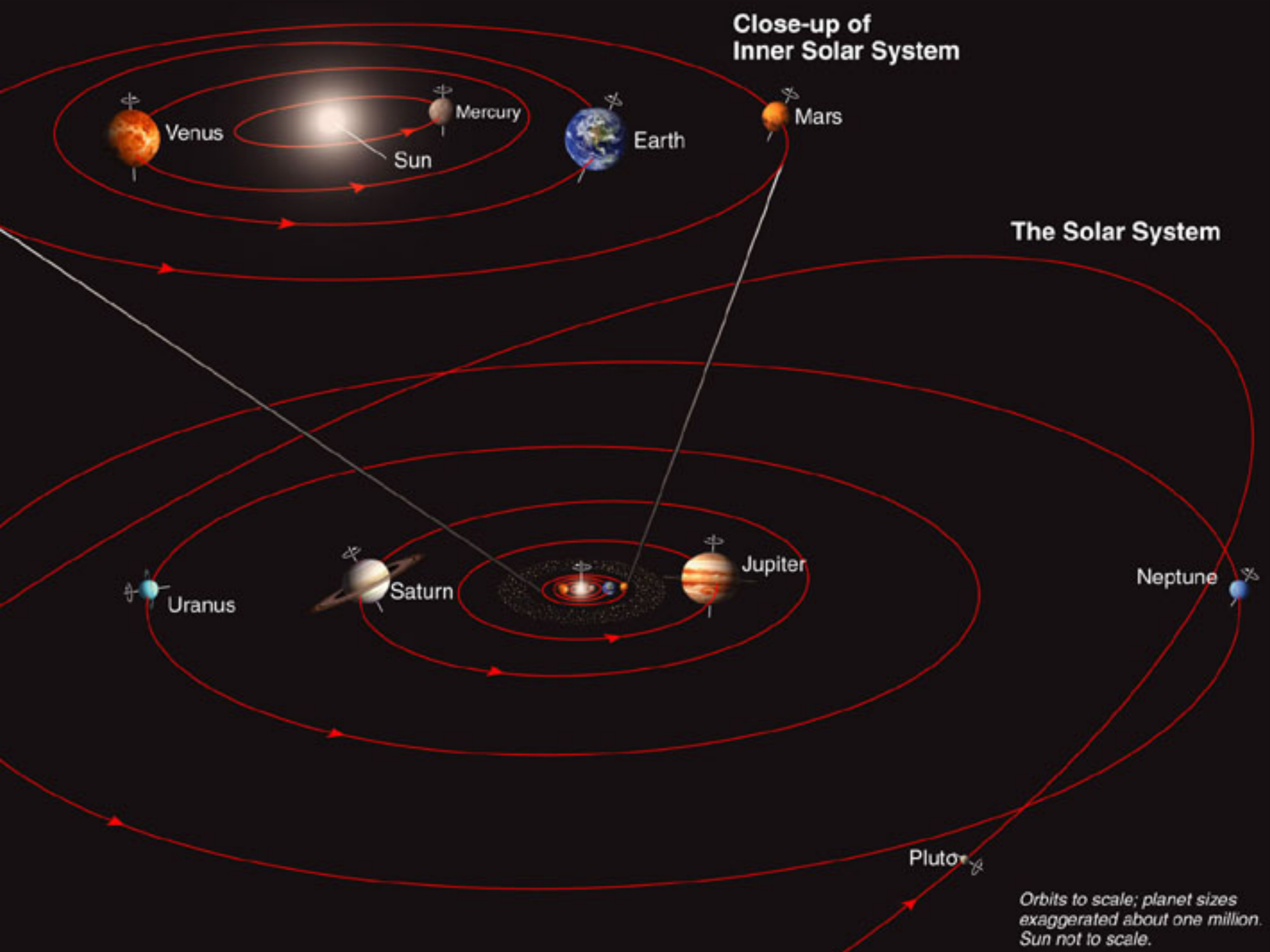
Ay 1 – Lecture 6



**Other Worlds:
Our Solar System
and the Others**

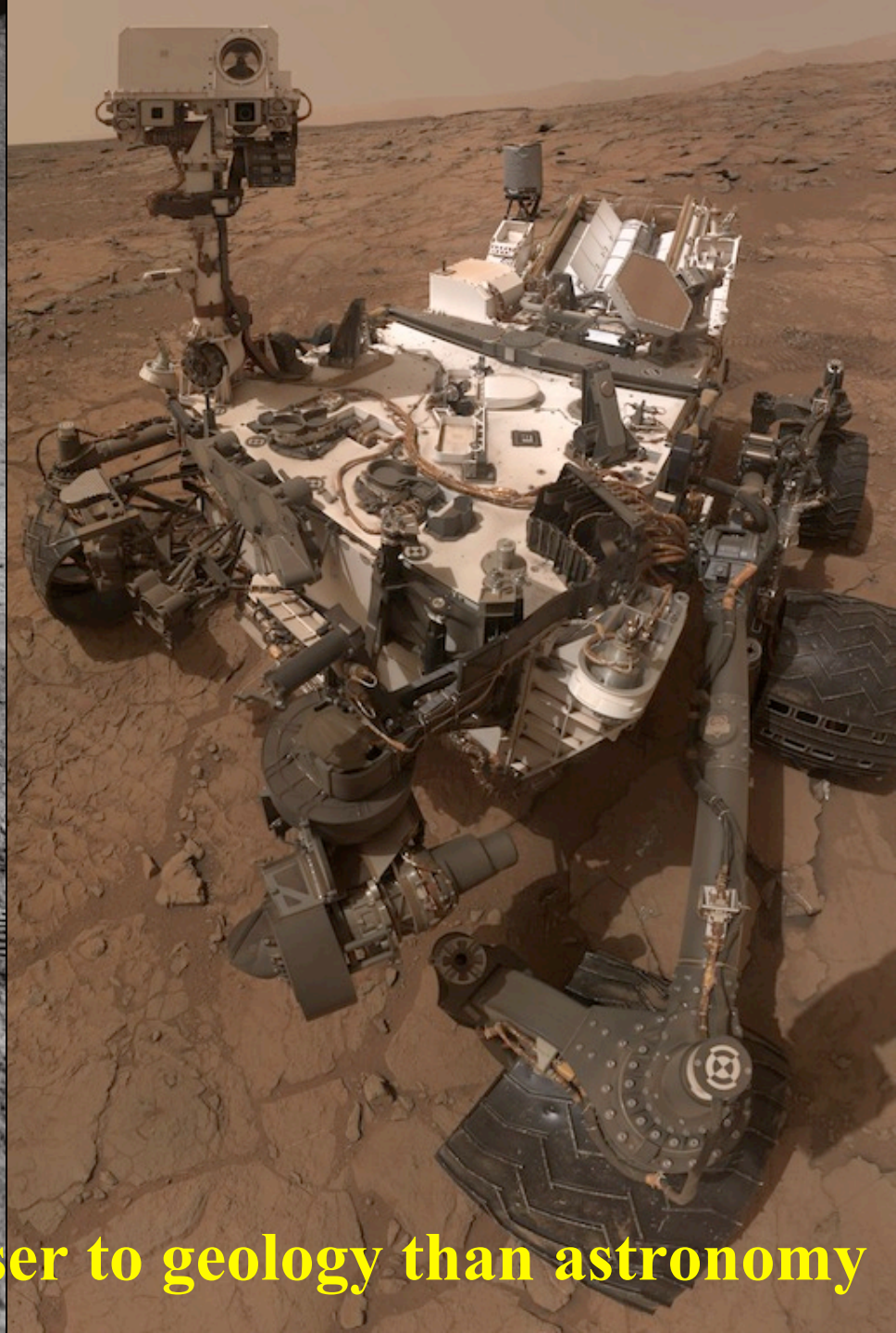
6.1 Contents of the Solar System





Planetary Demographics

Name	Distance from Sun (AU) ^[2]	Revolution Period (y)	Diameter (km)	Mass (10 ²³ kg)	Density (g/cm ³) ^[3]
Mercury	0.39	0.24	4,878	3.3	5.4
Venus	0.72	0.62	12,120	48.7	5.2
Earth	1.00	1.00	12,756	59.8	5.5
Mars	1.52	1.88	6,787	6.4	3.9
Jupiter	5.20	11.86	142,984	18,991	1.3
Saturn	9.54	29.46	120,536	5686	0.7
Uranus	19.18	84.07	51,118	866	1.3
Neptune	30.06	164.82	49,660	1030	1.6



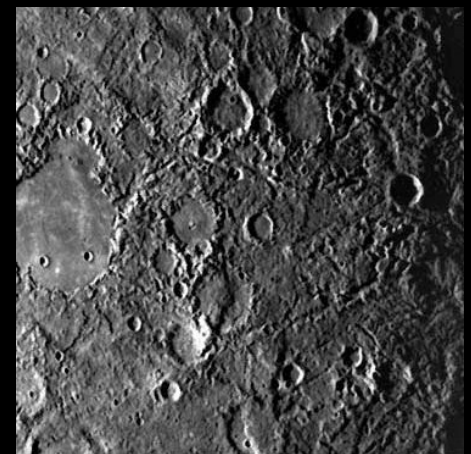
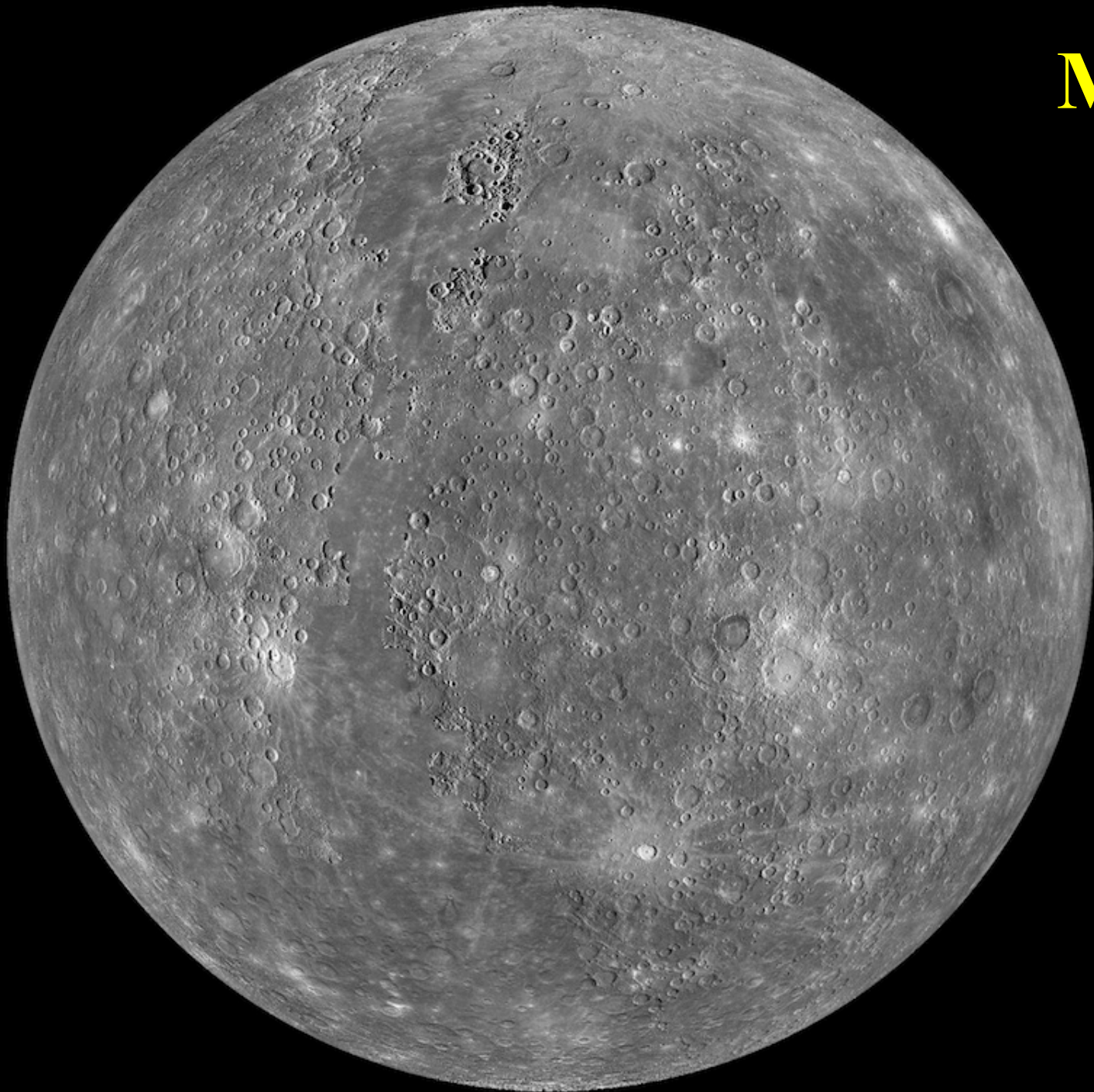
Planetary Science: closer to geology than astronomy

Three Kinds of Planets

- Rocky: inner Solar system, smaller, high density, composed of heavier elements
 - Mercury, Venus, Earth, Mars
- Gas giants: Outer Solar system, large, massive, lower densities, lighter elements are abundant
 - Jupiter, Saturn, Uranus, Neptune
- Dwarf planets: Very Outer Solar system, low mass, small, icy
 - Pluto, Sedna, Eris, Makemake, Ceres, etc.



Mercury

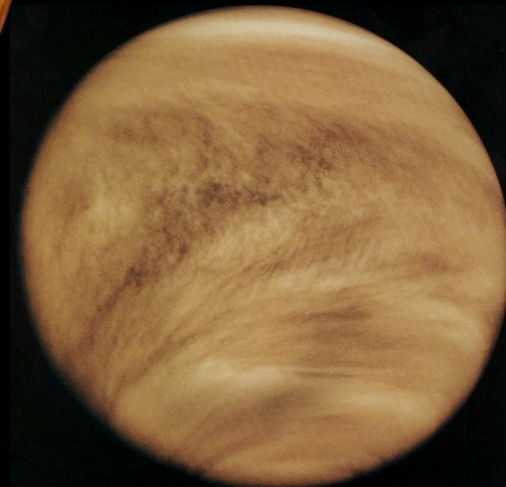


Venus

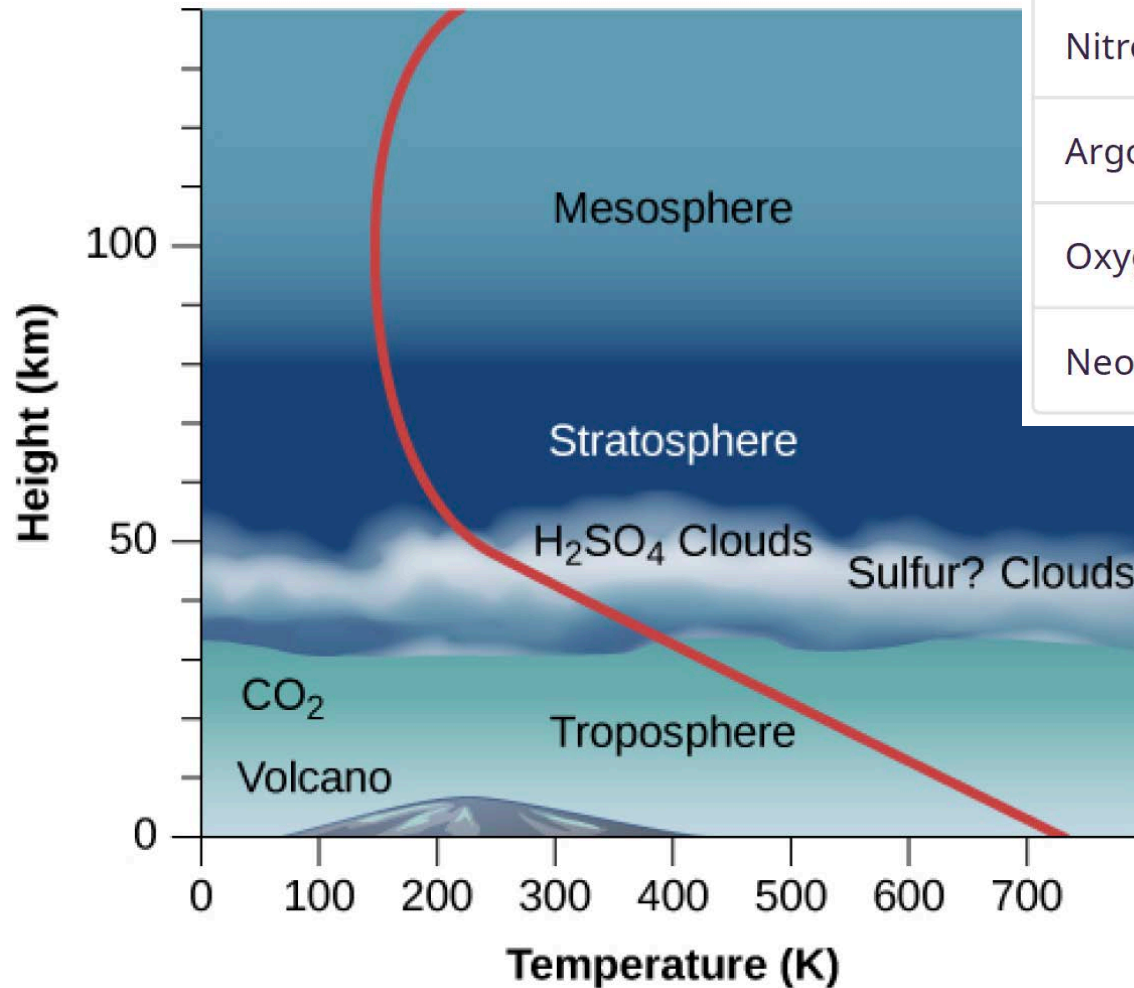
< Magellan Radar



Pioneer UV



Runaway Greenhouse Effect on Venus

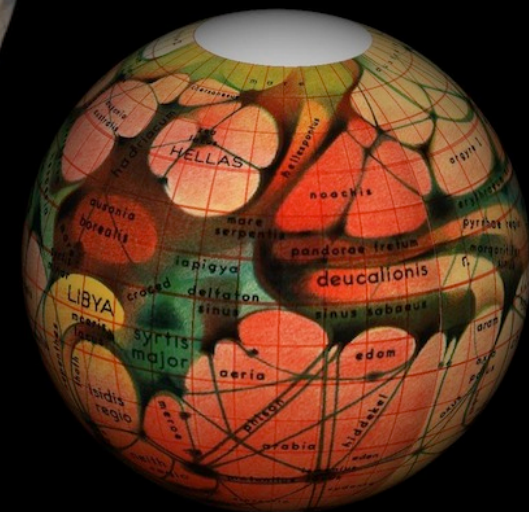
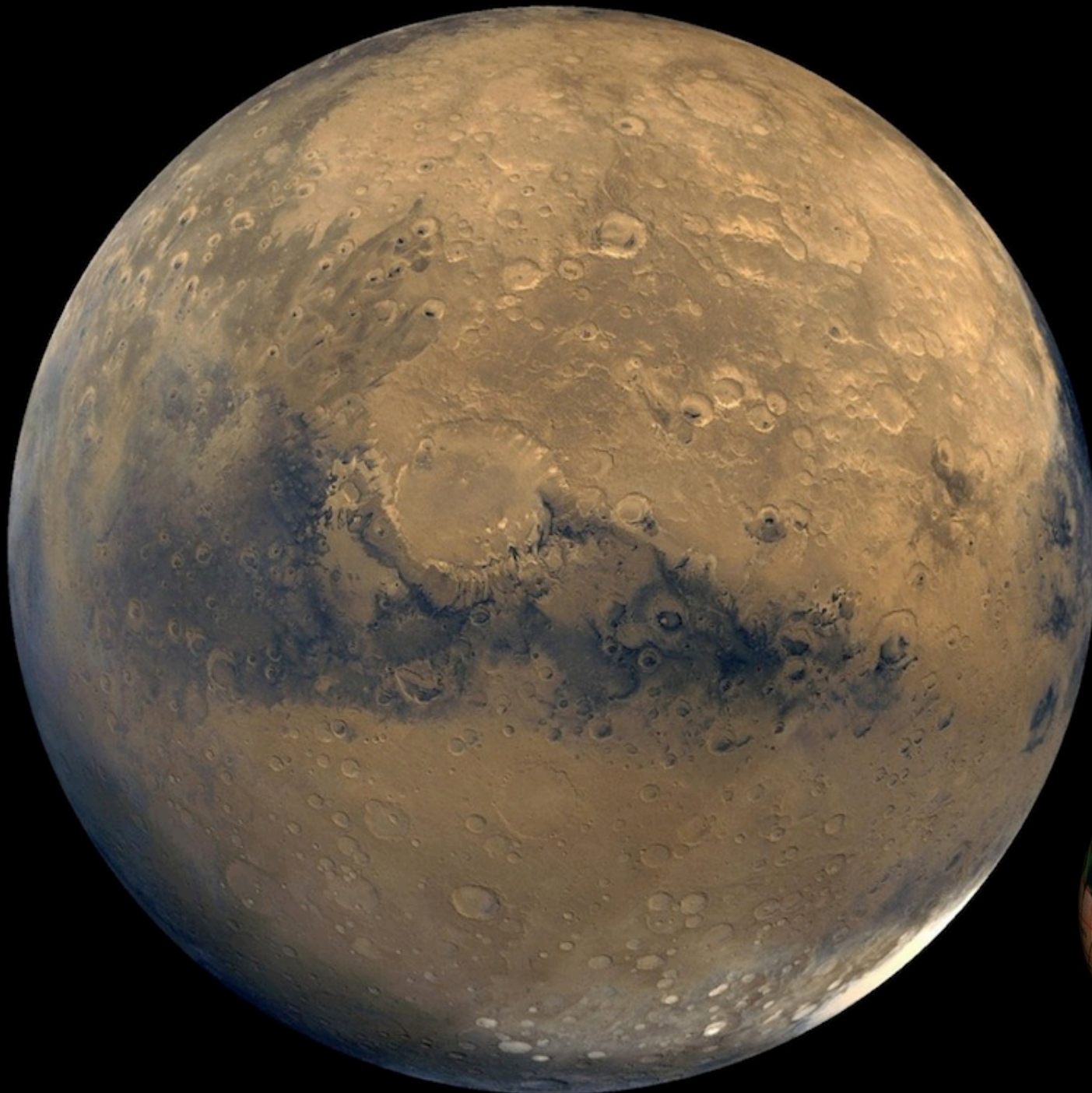


Gas	Earth	Venus
Carbon dioxide (CO ₂)	0.03%	96%
Nitrogen (N ₂)	78.1%	3.5%
Argon (Ar)	0.93%	0.006%
Oxygen (O ₂)	21.0%	0.003%
Neon (Ne)	0.002%	0.001%

Earth's Moon



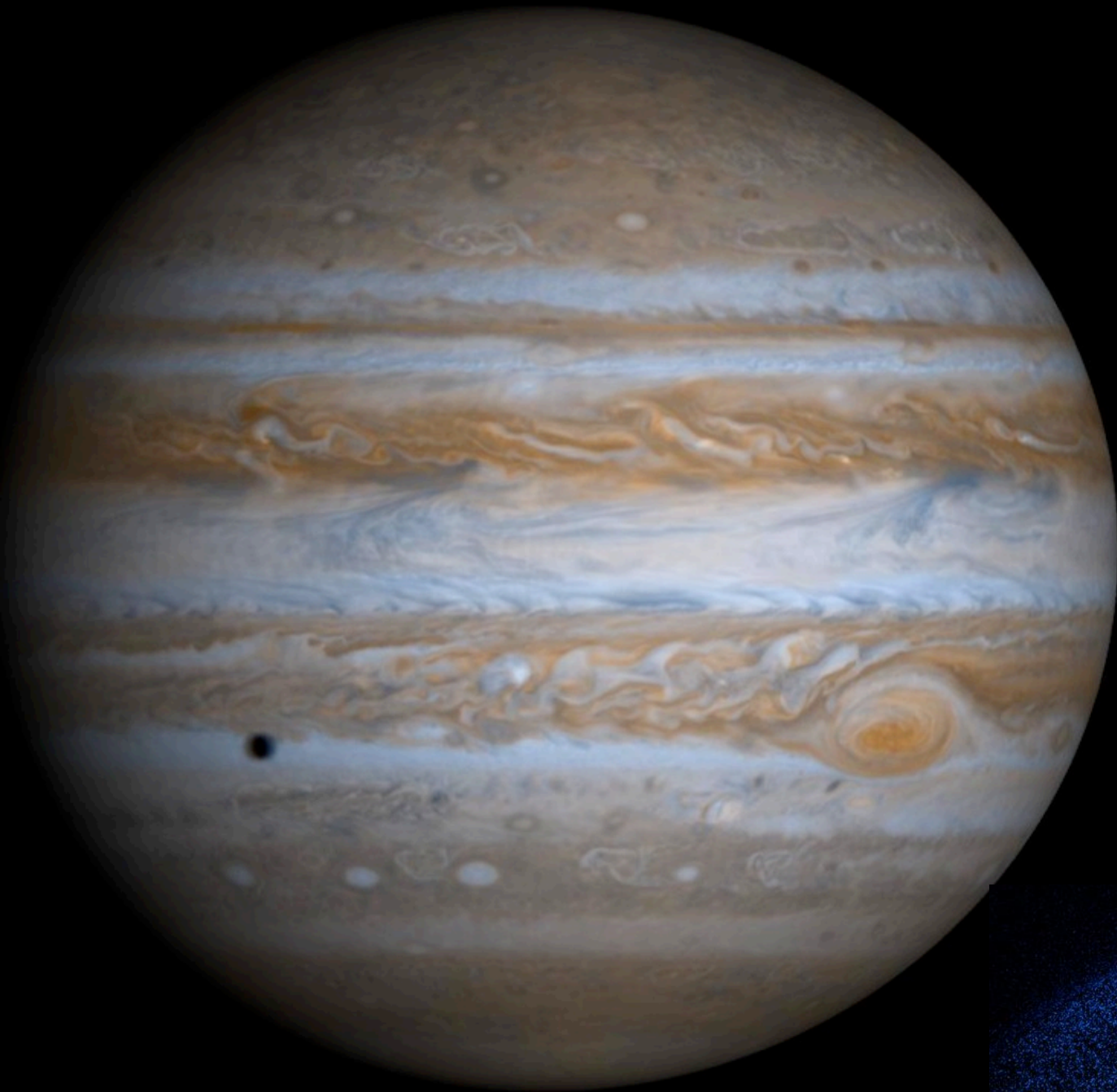
Mars

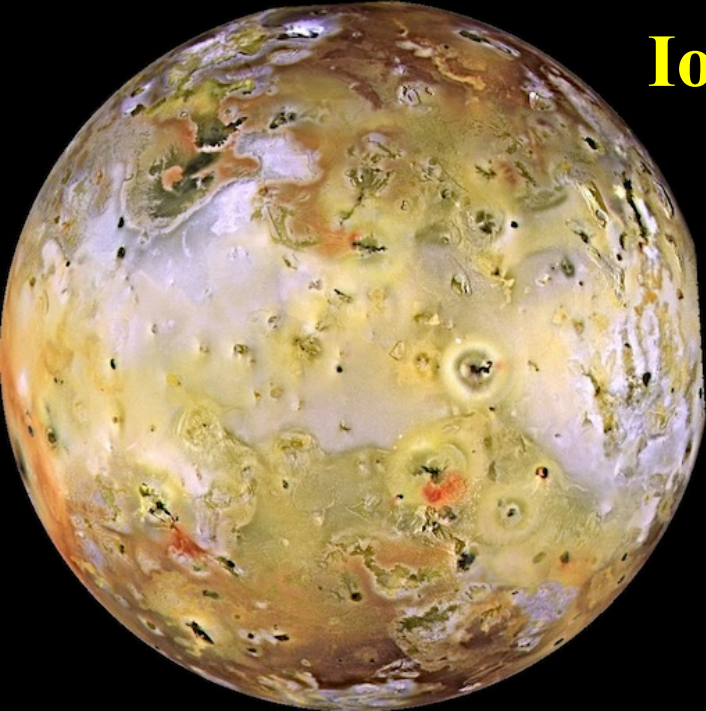




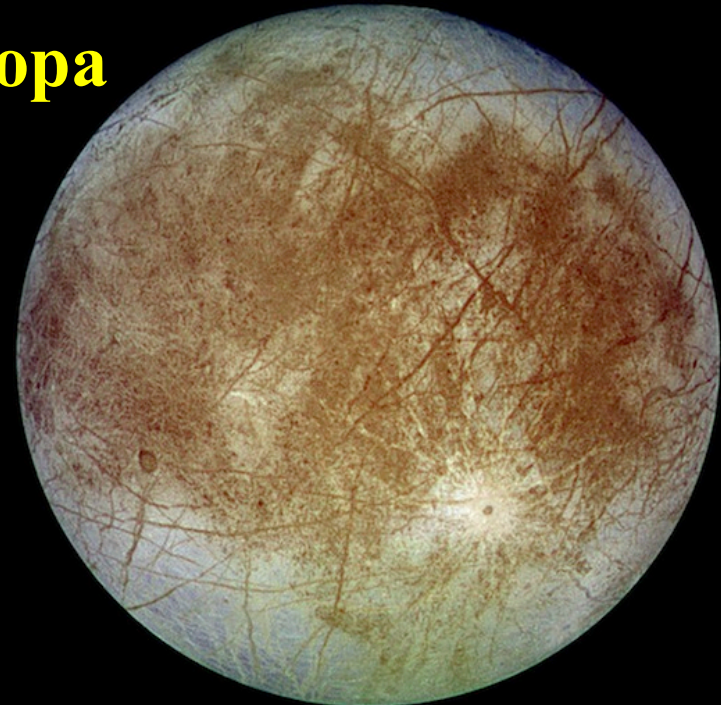
Water on Mars

Jupiter





Io

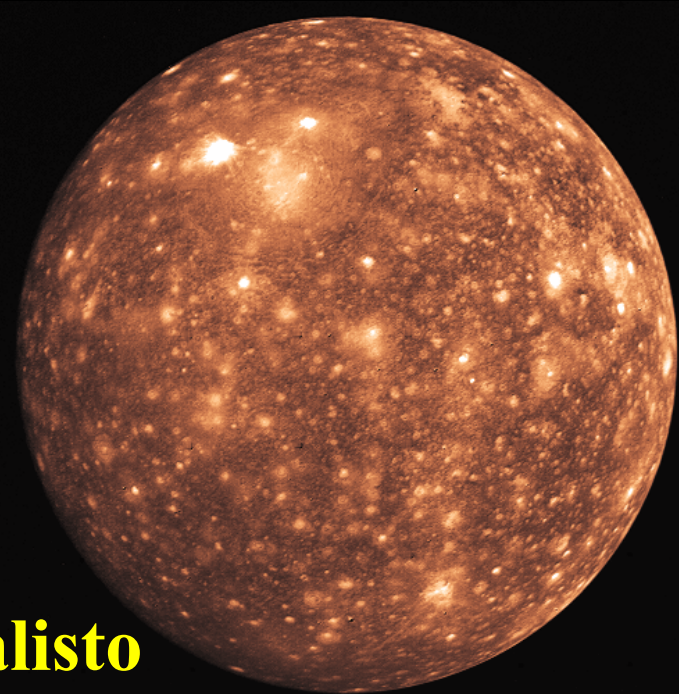


Europa

**Jupiter's
Moons**



Ganymede



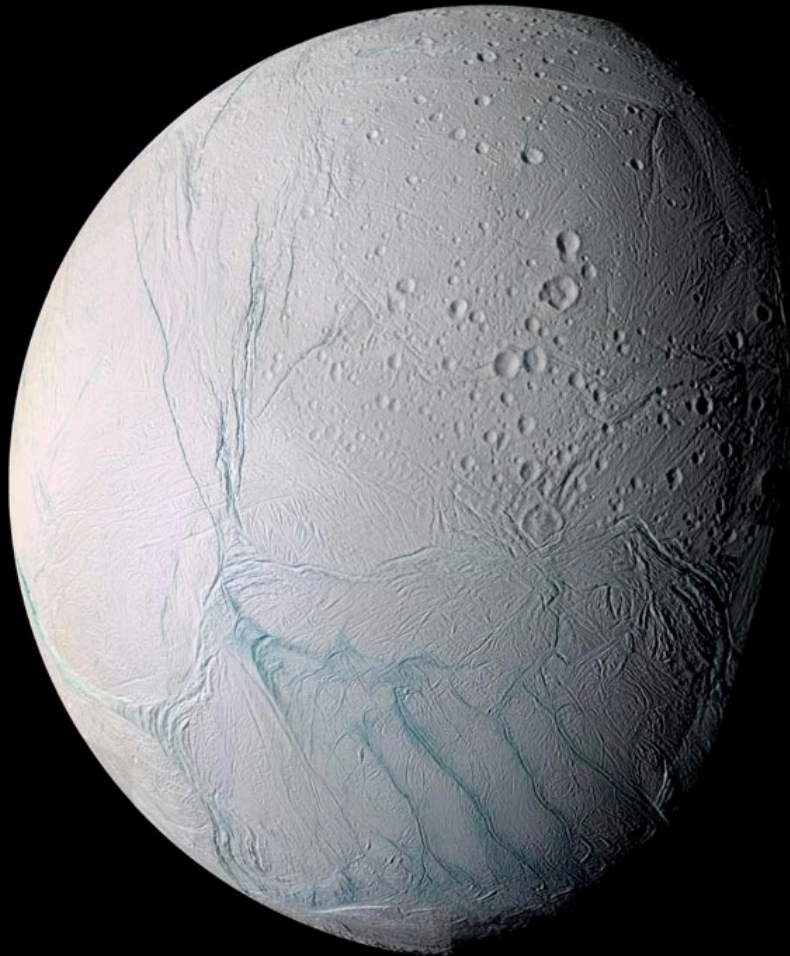
Calisto

Saturn

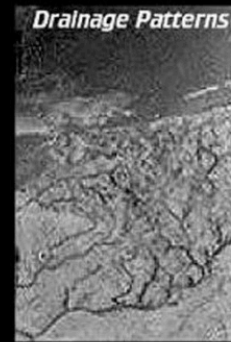


Saturn's Moons

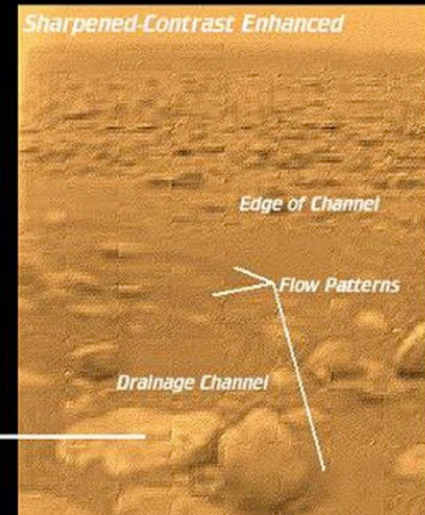
Enceladus



Titan



15 cm (6 inches)

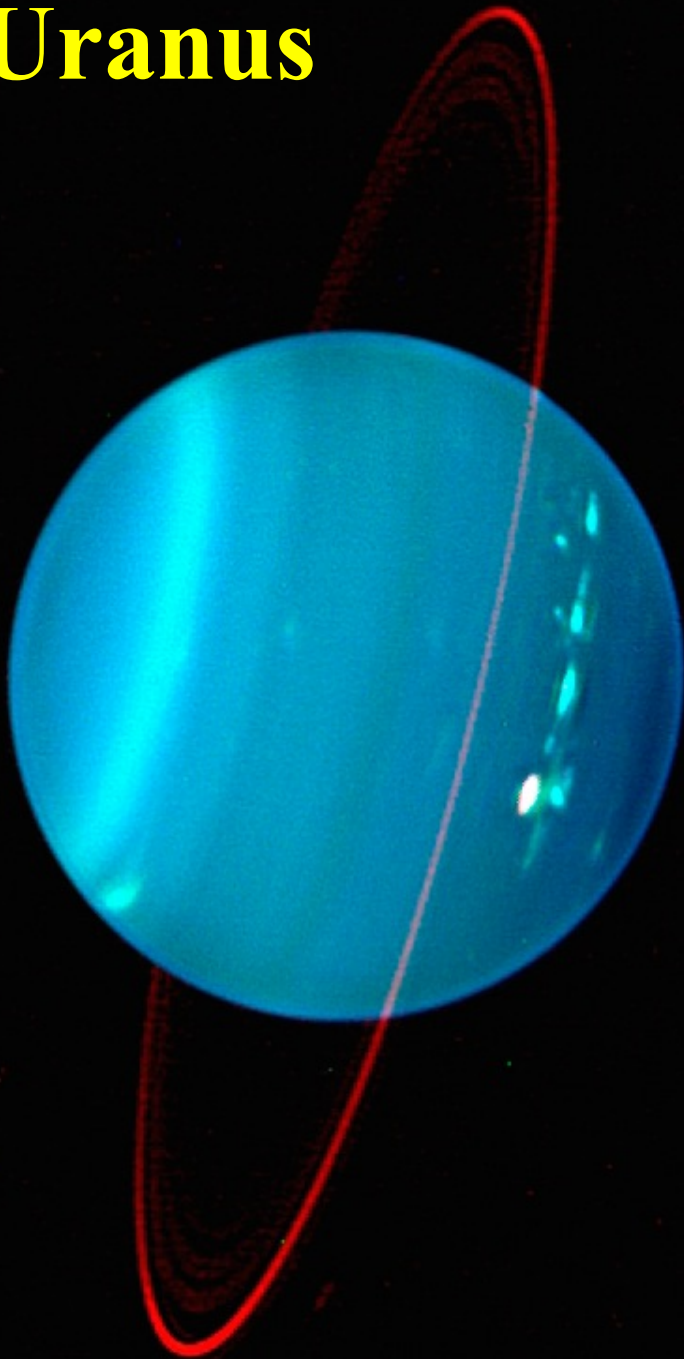


The image features two icy moons, Europa and Enceladus, positioned on either side of a central blue vertical band. Europa, on the left, is shown in a light tan and white color palette, with a complex network of reddish-brown linear features across its surface. Enceladus, on the right, is depicted in shades of blue and grey, also showing a network of linear features. The background is a gradient of blue, suggesting a sky or space environment.

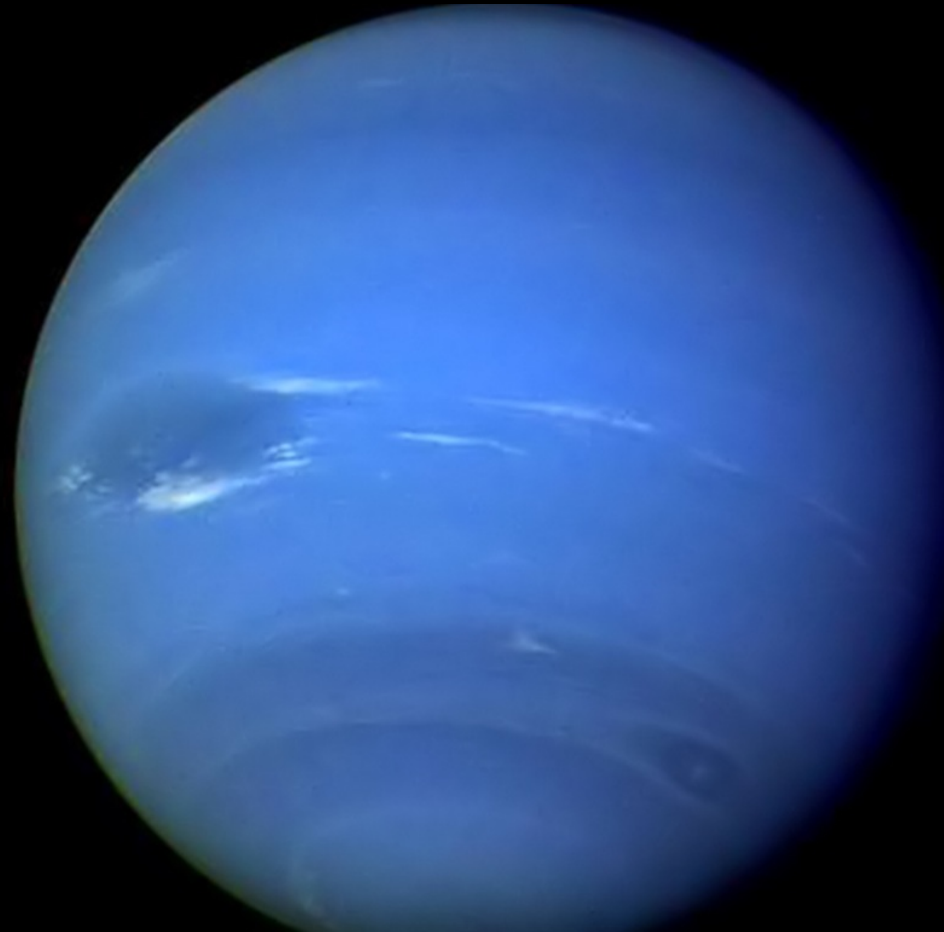
Oceans under the ice crust

on Europa
and Enceladus

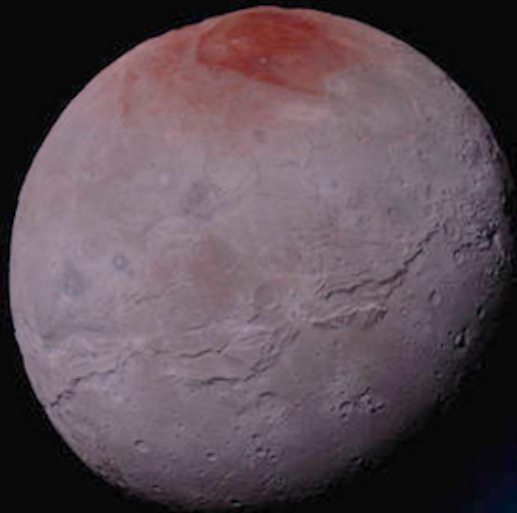
Uranus



Neptune



Pluto and Charon



Pluto Killer
(Mike Brown)

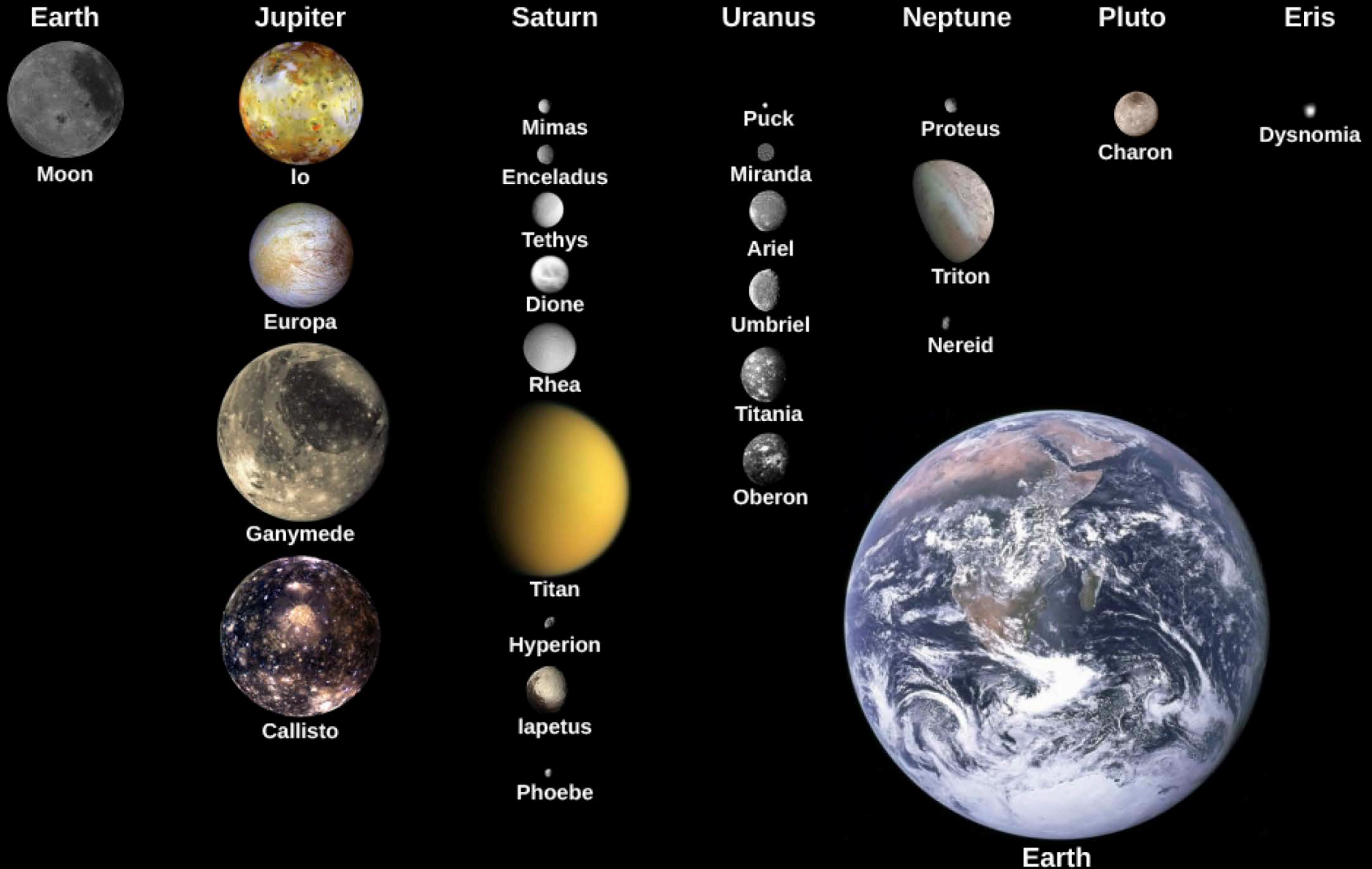


Planet Nine (?)



Predicted by Konstantin Batygin and
Mike Brown (but not yet discovered)

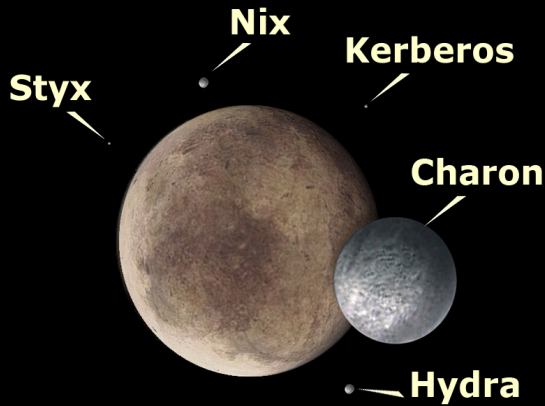
Larger Moons in the Solar System



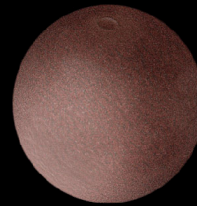
Largest known trans-Neptunian objects (TNOs)



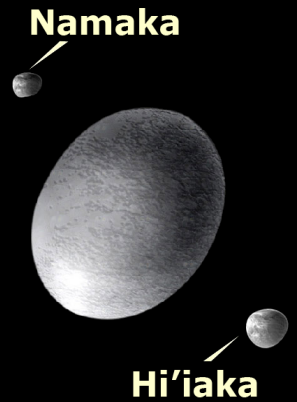
Eris



Pluto



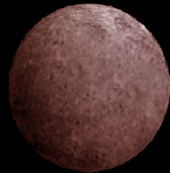
Makemake



Haumea



Sedna



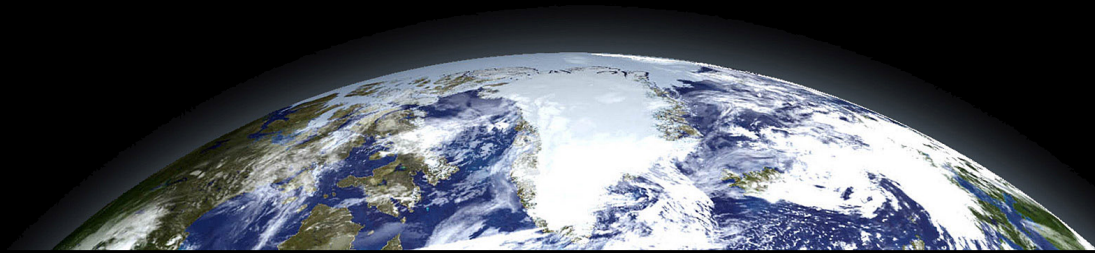
2007 OR₁₀



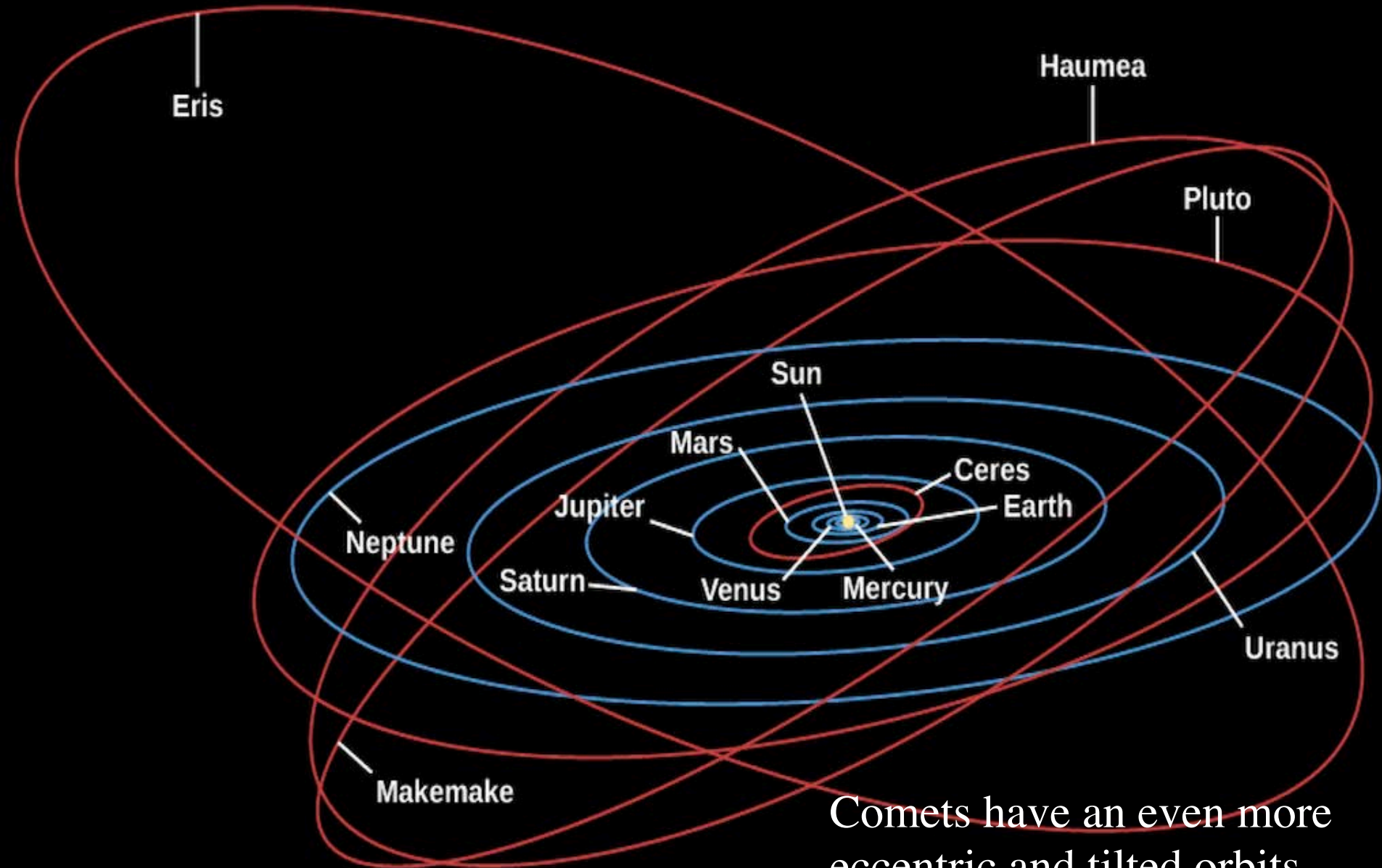
Quaoar



Orcus



Orbits in the Solar System



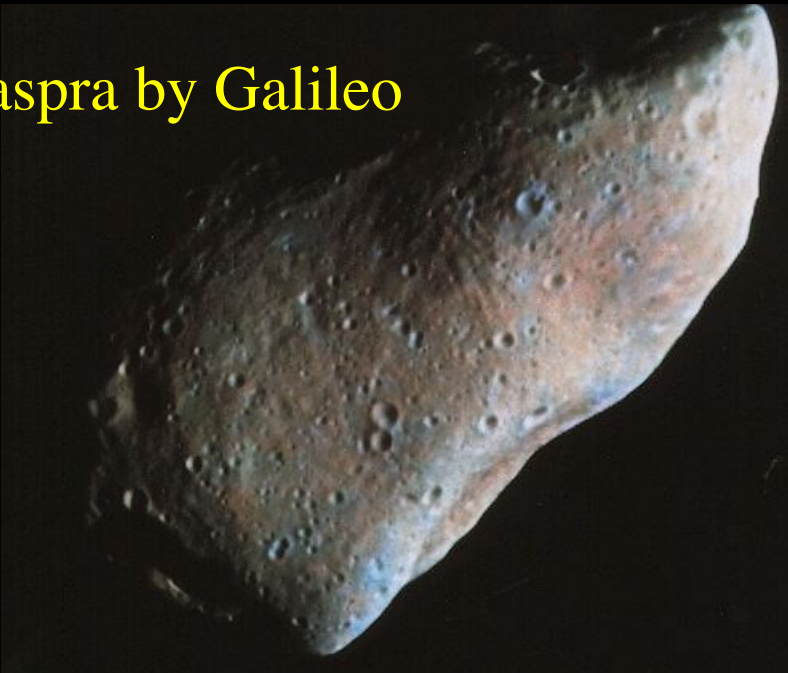
Comets have an even more eccentric and tilted orbits

Asteroids: Leftover Rocky Planetesimals



Mathilda by NEAR

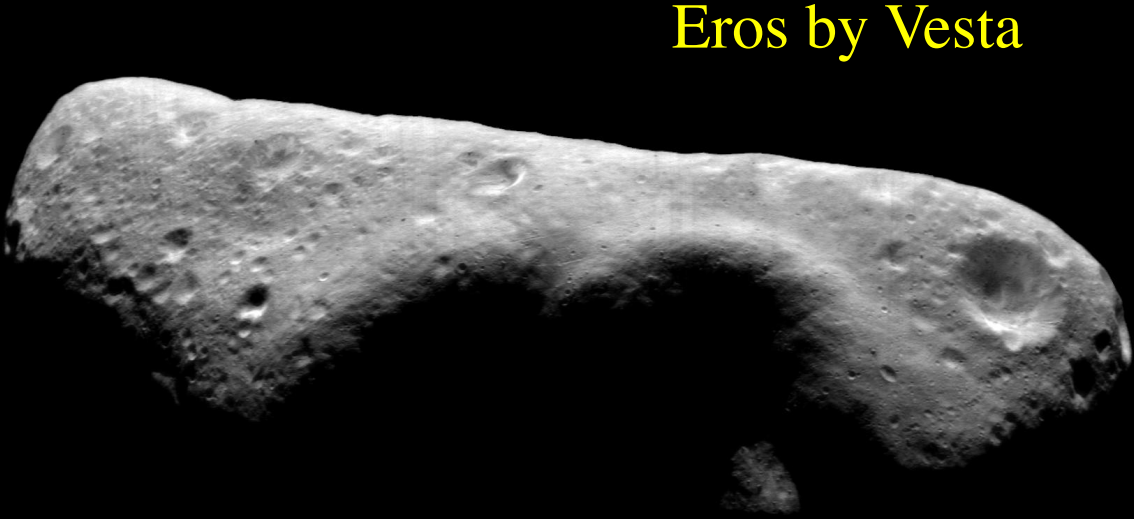
30 km



Gaspra by Galileo



Ida by Galileo



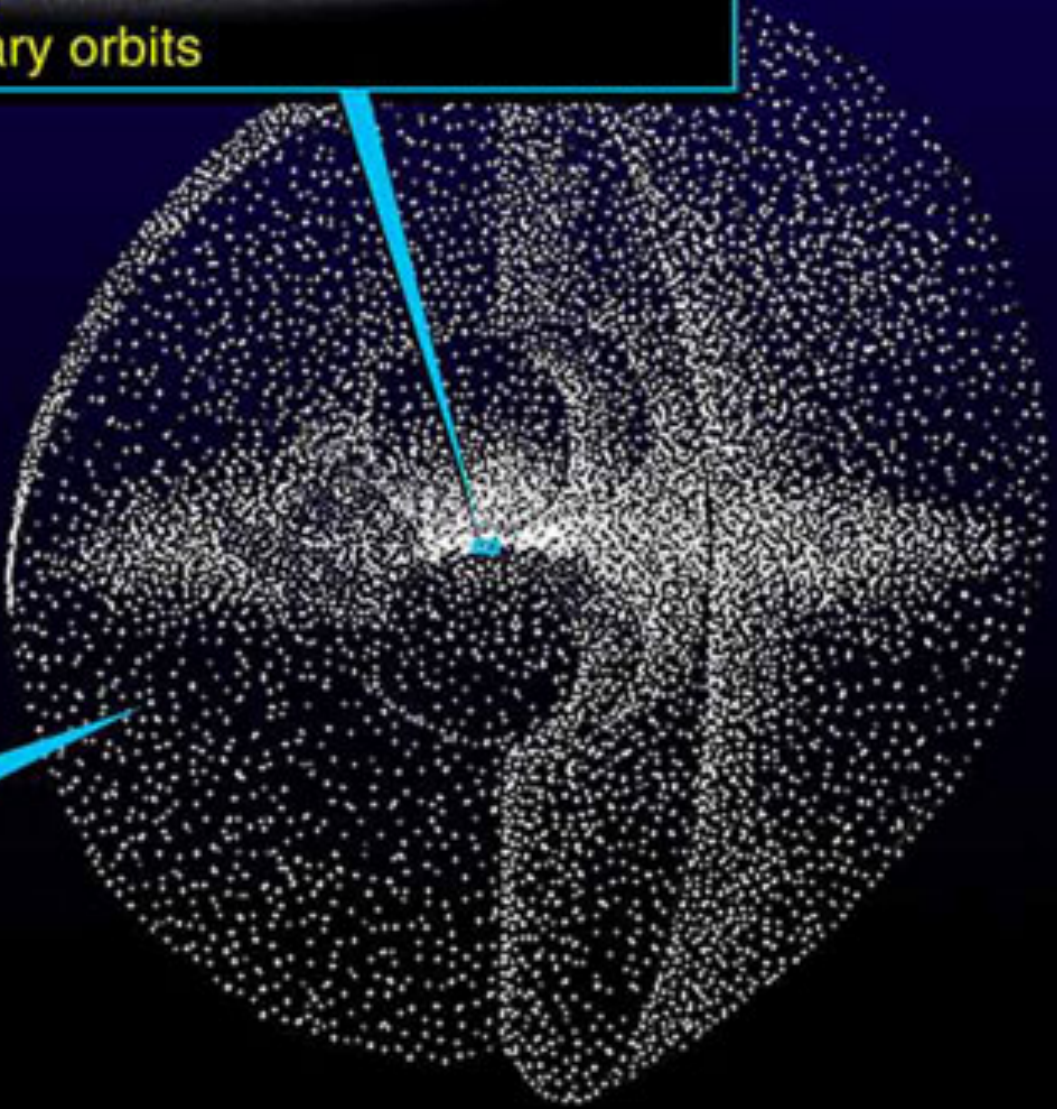
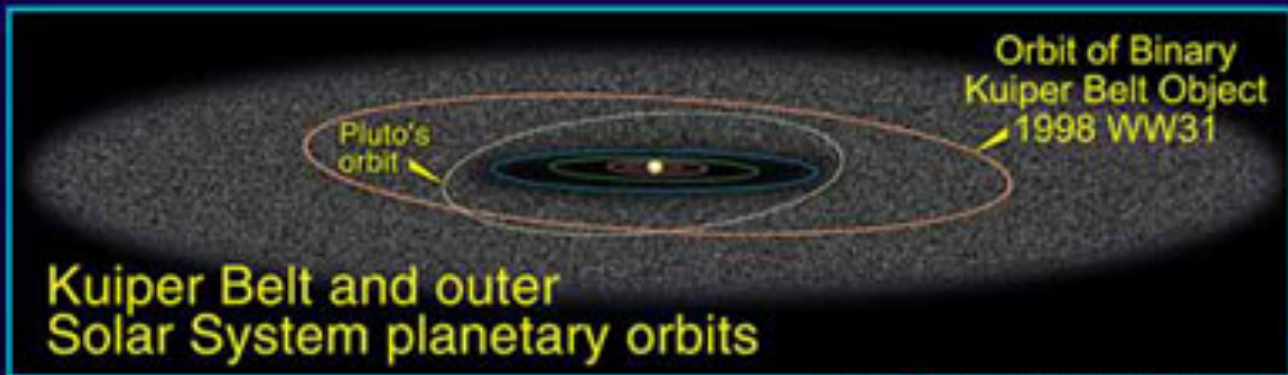
Eros by Vesta

Comets: Leftover Icy Planetesimals

Hartle 2



Halley



The Oort Cloud (comprising many billions of comets)

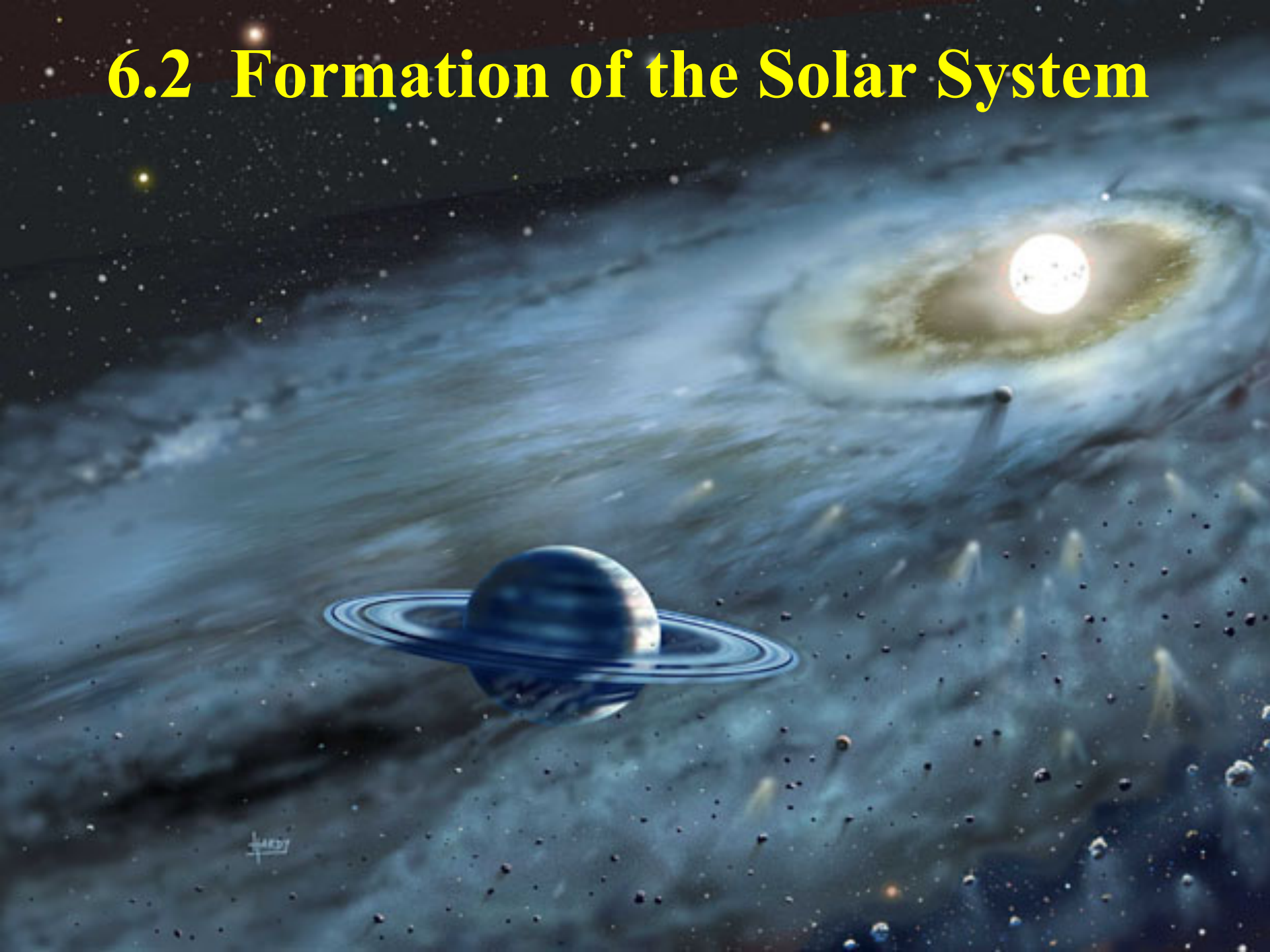
Oort Cloud cutaway drawing adapted from Donald K. Yeoman's illustration (NASA, JPL)

Zodiacal Dust: Leftover Protoplanetary Disk Dust





6.2 Formation of the Solar System



The Idea of Planetesimals and the Origin of the Solar System

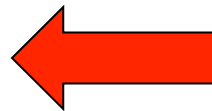
Everywhere in the solar nebula, tiny pieces of matter started condensing from the gas



At different places in the solar nebula, these “little bits of grit” were different compounds

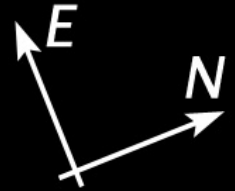


Eventually, these planetesimals collected into objects the size of planets. Gravity got into the act when the planetesimals got big



These small pieces of matter stuck to others, making larger sized blocks (the **planetesimals**)

Fomalhaut HST ACS/HRC



Dust ring

No data

Scattered
starlight
"noise"

Location of
Fomalhaut

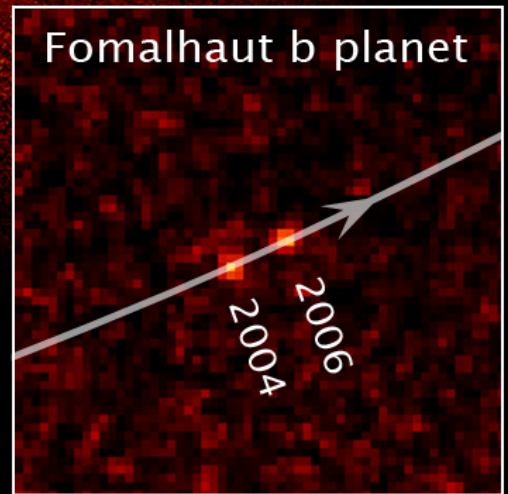
Coronagraph
mask

Large
planetesimals
have probably
already formed
in here

No data

Background Star

100 AU 13"

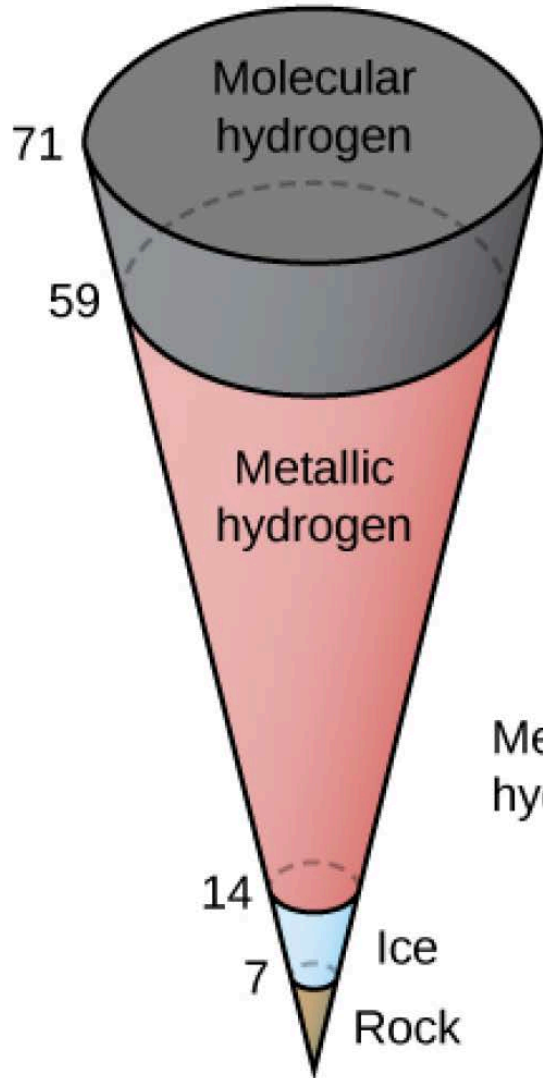


Masses and Compositions of the Major Planets

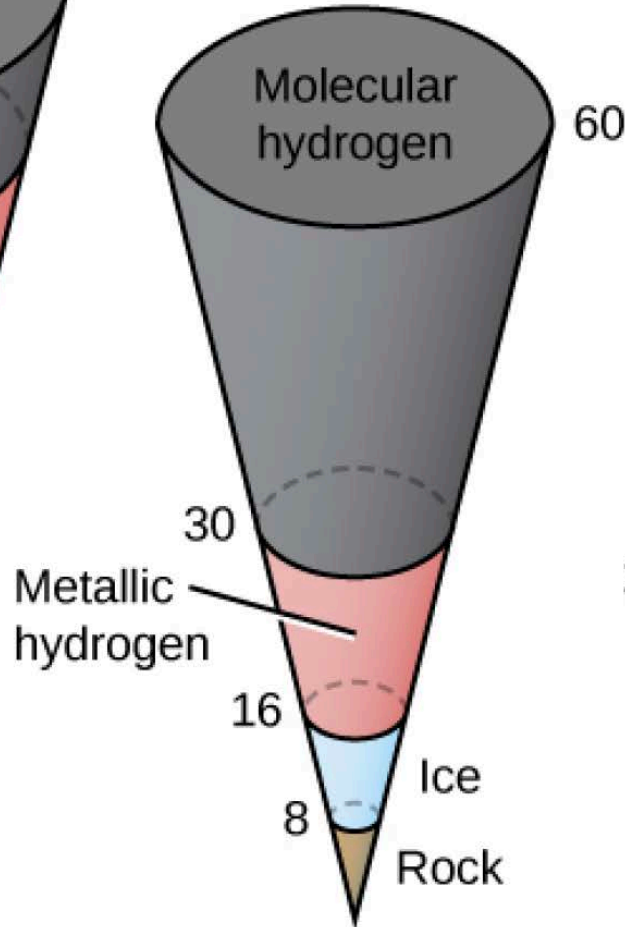
- At the location of the terrestrial planets, there was not much mass in the planetesimals, since they were formed of heavier, non-abundant elements
- In the outer solar system, there was more mass in the planetesimals, since they were formed of abundant, hydrogen-bearing compounds. Apparently, they produced more massive planetesimals that incorporated the hydrogen and helium gas that makes up most of Jupiter and Saturn
- At the position of the Earth, only silicates and other more “refractory” substances would have precipitated from the vapor state. At Jupiter and beyond, ices of water, ammonia, methane, would have condensed

Composition of the Gas Giant Planets

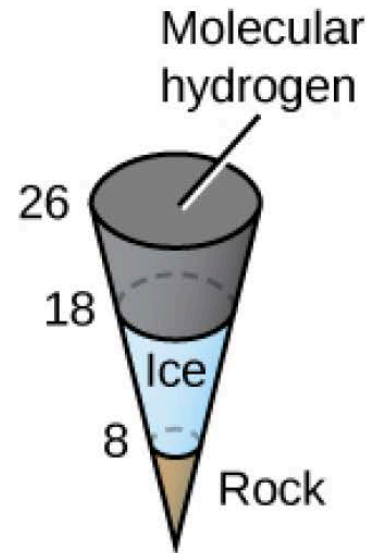
Jupiter



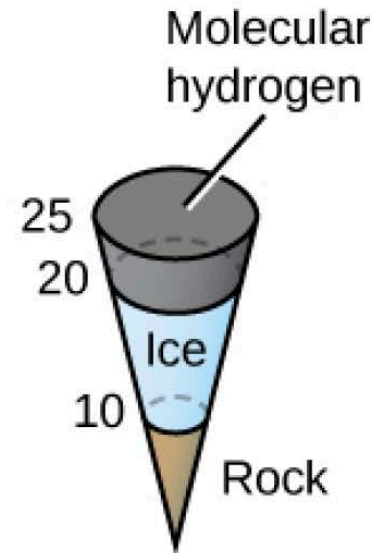
Saturn



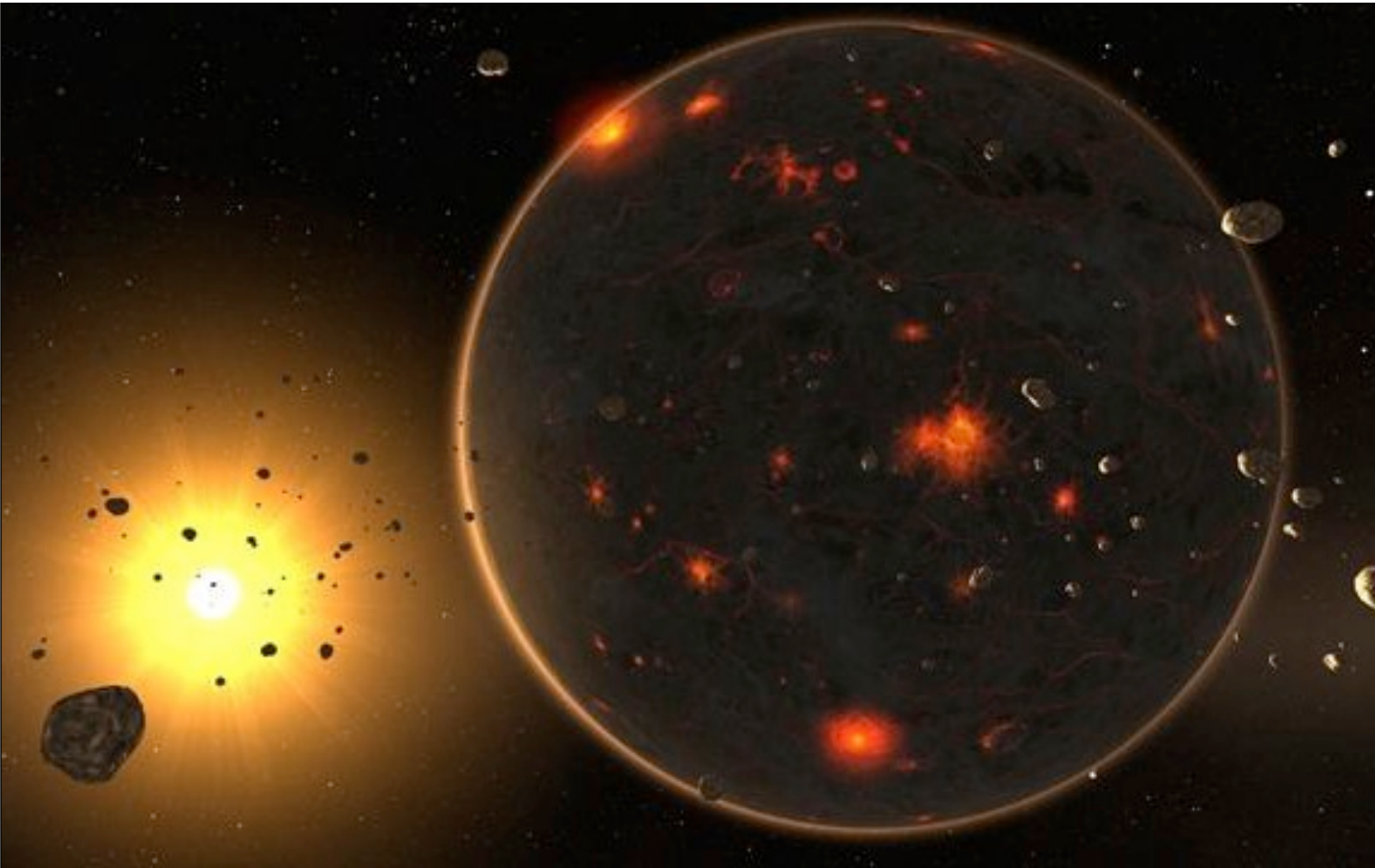
Uranus



Neptune



Late Heavy Bombardment

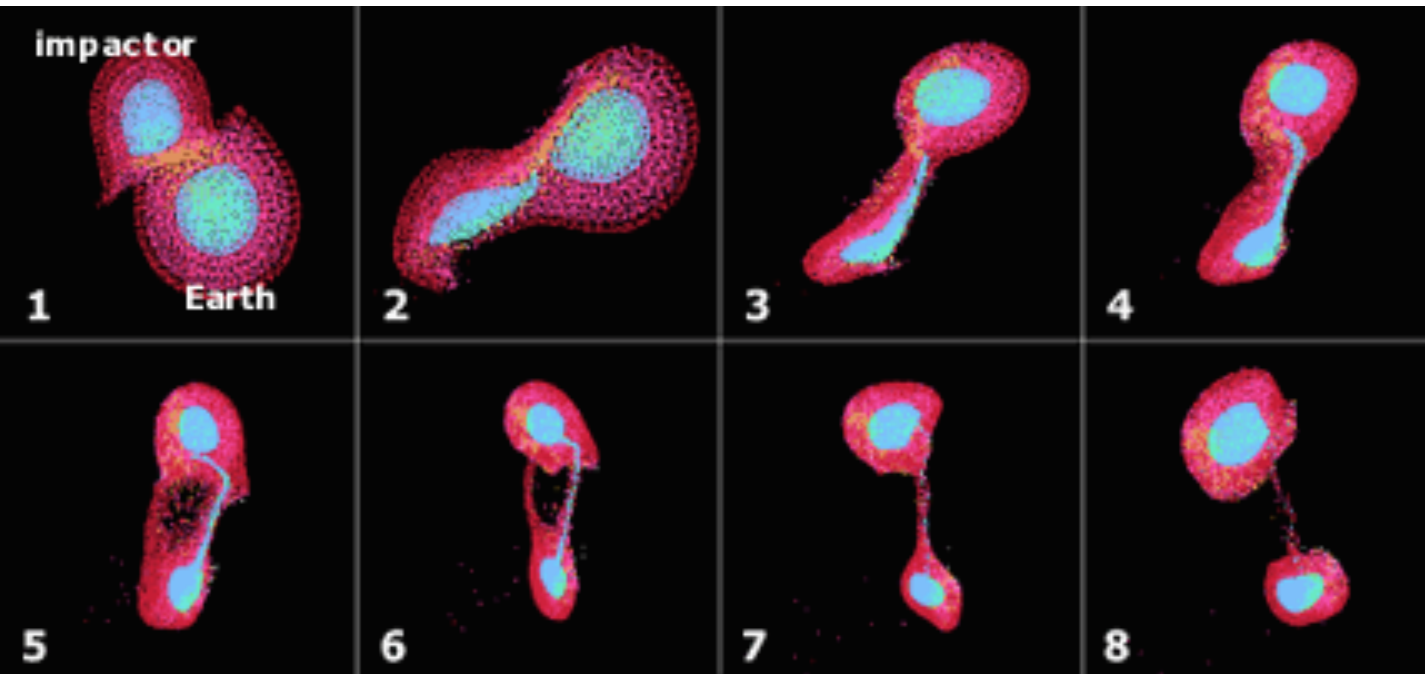


The Origin of the Moon

A Mars-sized protoplanet colliding with the proto-Earth



Moon condenses from the debris



Explains:

- Lunar composition
- Tilt of the Earth's axis

(Courtesy of A. G. W. Cameron, Harvard College Observatory.)

Cretaceous-Tertiary Impact Extinction



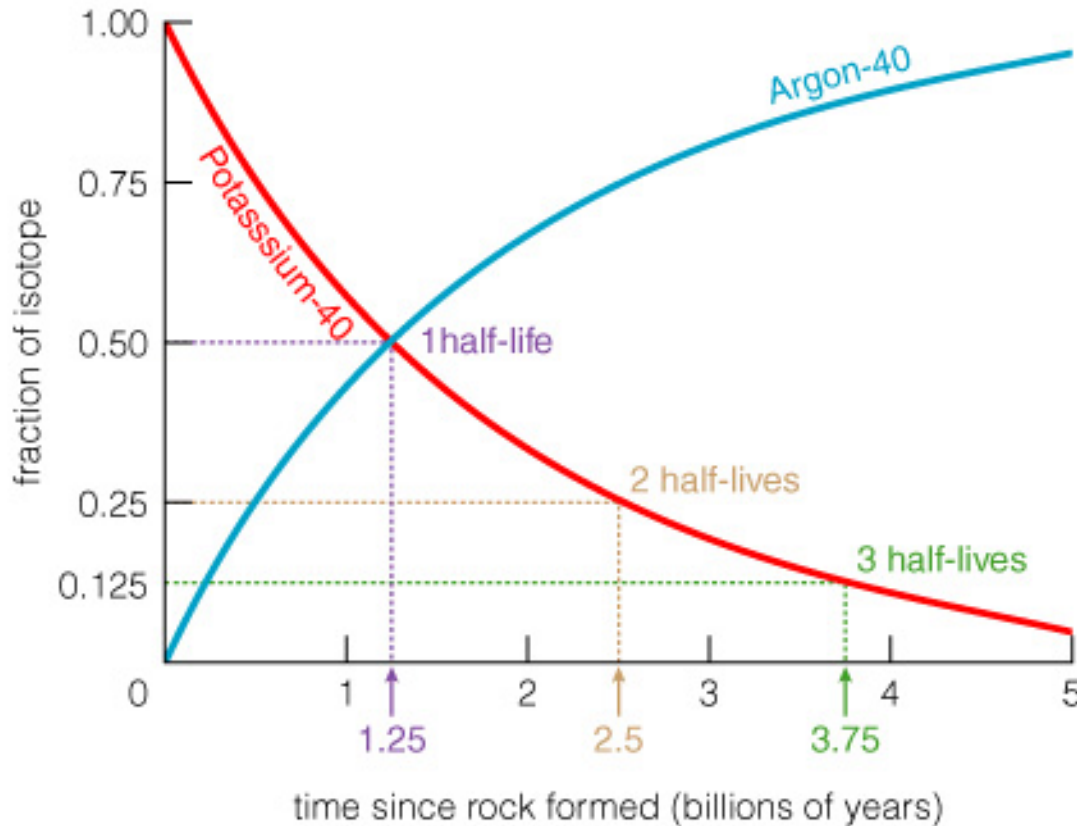
The Impacts Continue

Tunguska >

Large meteor crater, Arizona

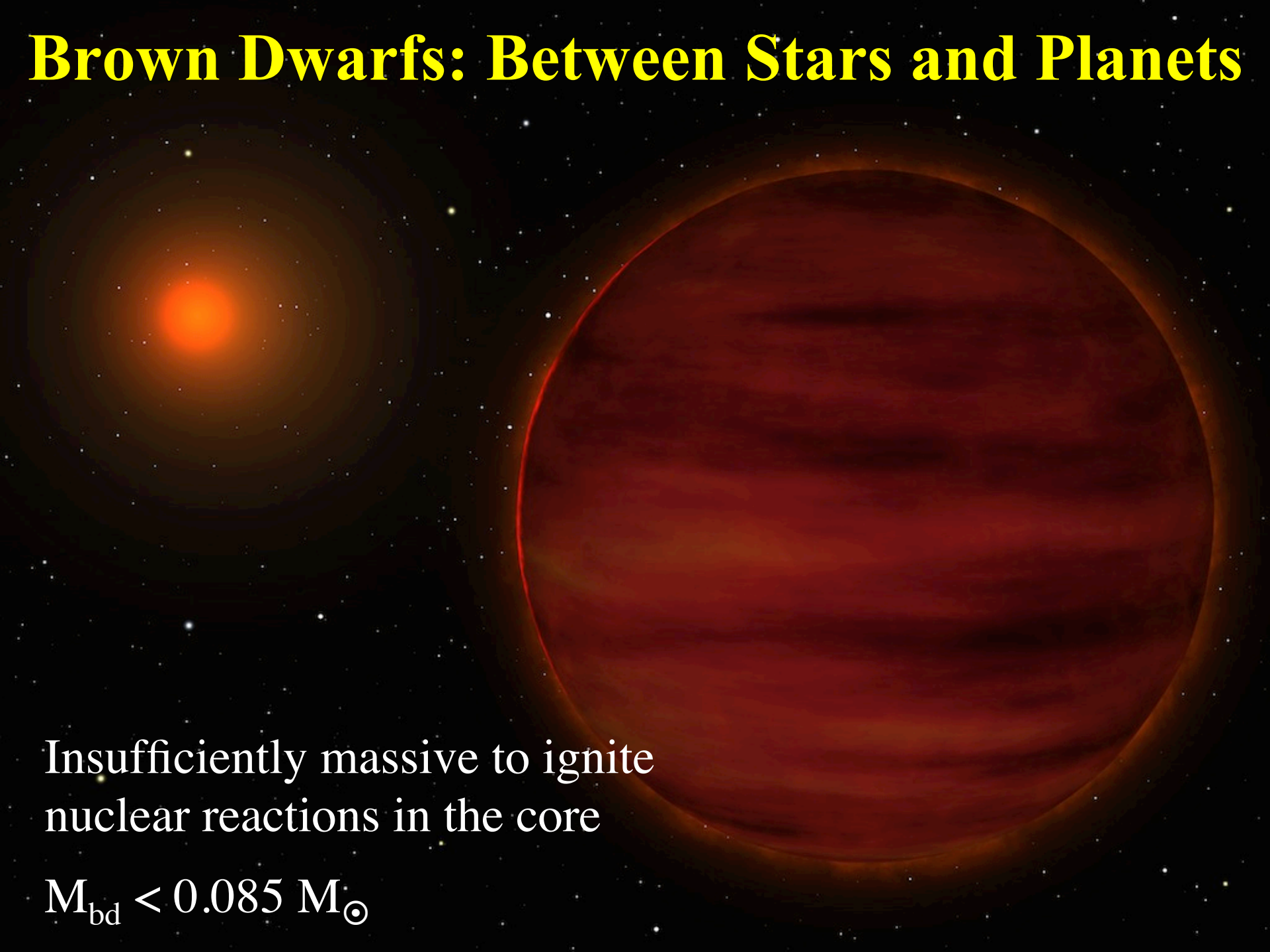


When Did the Planets Form?



- Some isotopes decay into other nuclei
 - A **half-life** is the time for half the nuclei in a substance to decay
 - Relative abundances of these isotopes then give us the age
-
- Radiometric dating tells us that oldest moon rocks are 4.4 billion years old
 - Oldest meteorites are 4.55 billion years old
 - Planets probably formed ~ 4.6 billion years ago

Brown Dwarfs: Between Stars and Planets



Insufficiently massive to ignite
nuclear reactions in the core

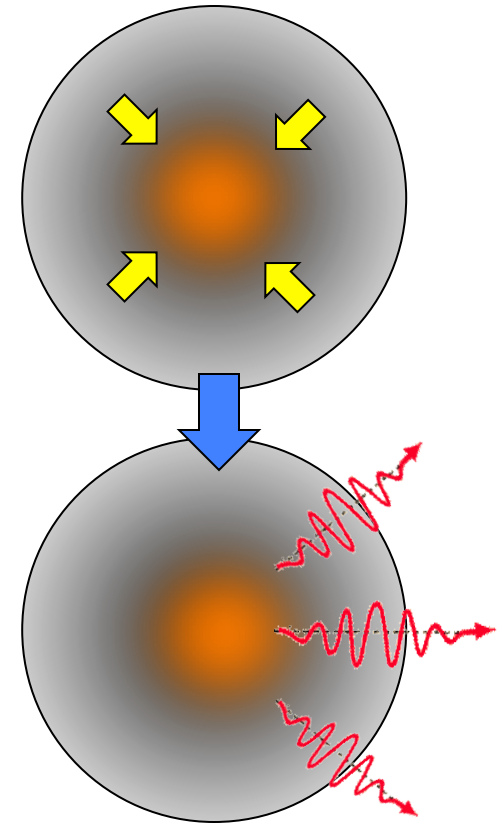
$$M_{\text{bd}} < 0.085 M_{\odot}$$

The Kelvin-Helmholtz Mechanism

As a planet cools, it shrinks

The release of the binding energy produces heat, that radiates away

For example, Jupiter, and all brown dwarfs



Total binding energy available divided by the luminosity gives the *Kelvin-Helmholtz time scale*
For Sun, that is ~ 18 million years



6.3 Planetary Atmospheres

The background image is a high-resolution view of a planetary atmosphere, showing intricate cloud patterns and a prominent dark band. The colors range from light beige and tan to deep red and brown, indicating different atmospheric layers and chemical compositions. The patterns are swirling and turbulent, characteristic of a gas giant's atmosphere.

How do you obtain an atmosphere?

- Gain volatiles by comet impacts
- Outgassing during differentiation
- Ongoing outgassing by volcanoes

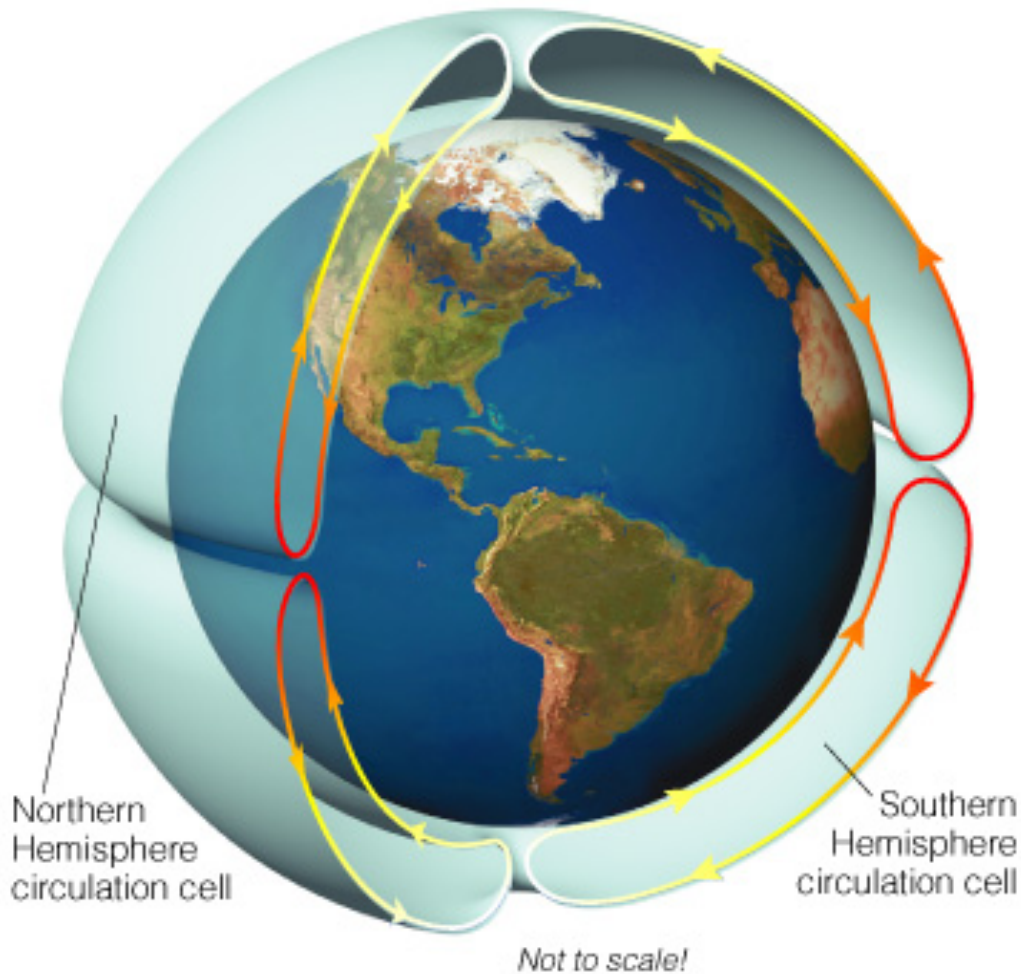


Keeping an Atmosphere

Atmosphere is *kept* by the world's gravity

- Low mass worlds = low gravity = almost no atmosphere
- High mass worlds = high gravity = thick atmosphere

Why are the winds blowing? The answer, my friends, is...

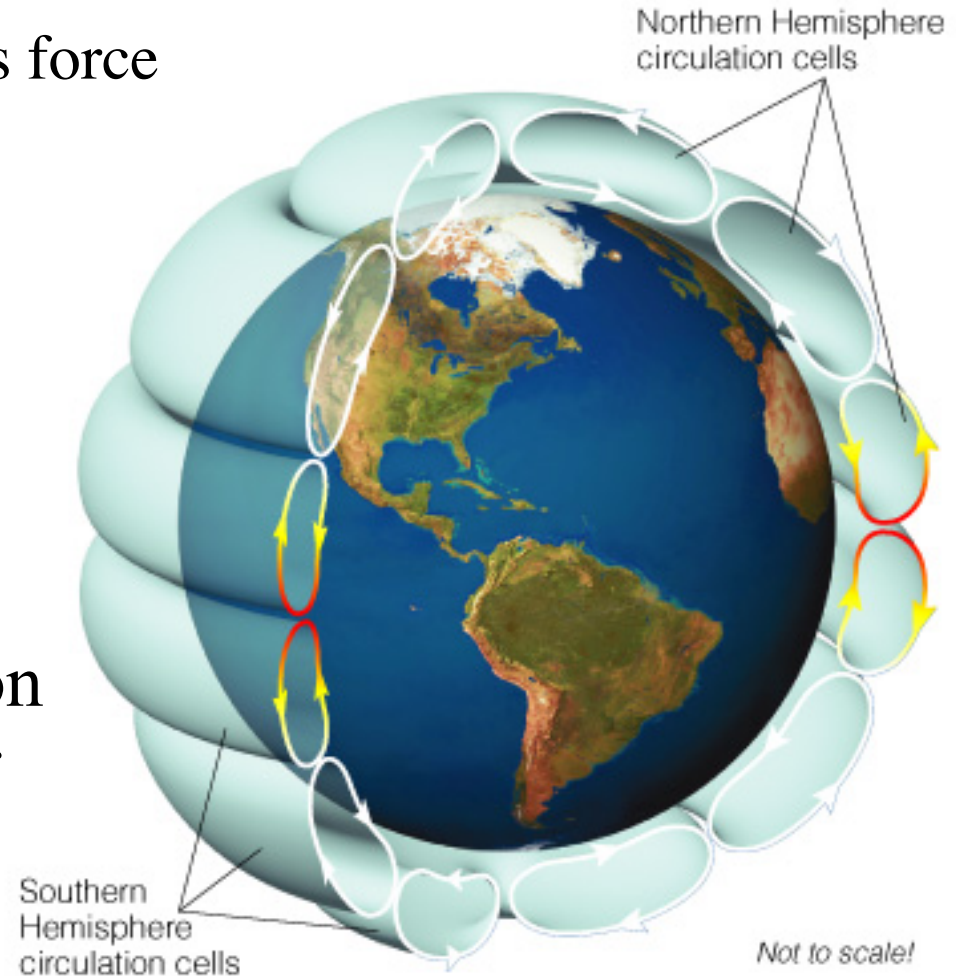
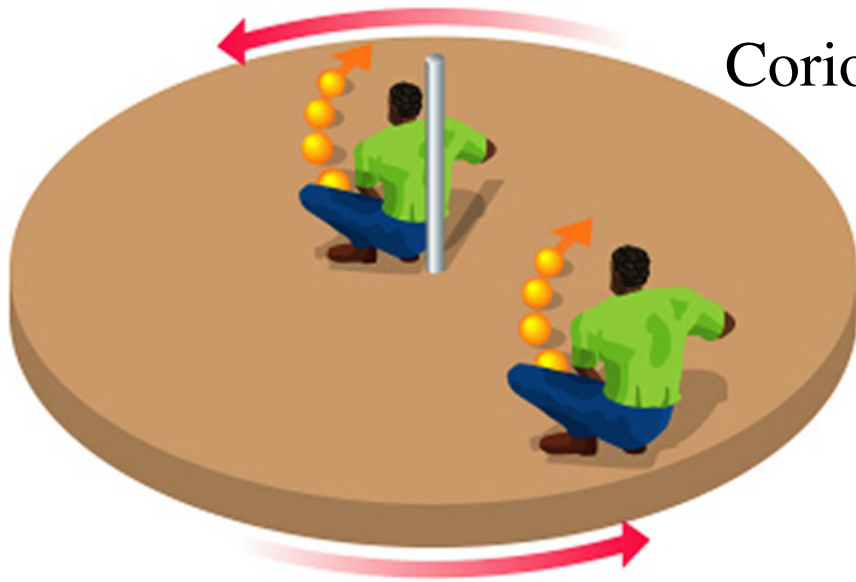


- Heated air rises at equator

←←←← Maximum
←←←← Sun warming

- Cooler air descends at poles

The planetary rotation also plays a role:



- On Earth the large circulation cell breaks up into 3 smaller ones, moving diagonally
- Other worlds have more or fewer circulation cells depending on their rotation rate

